

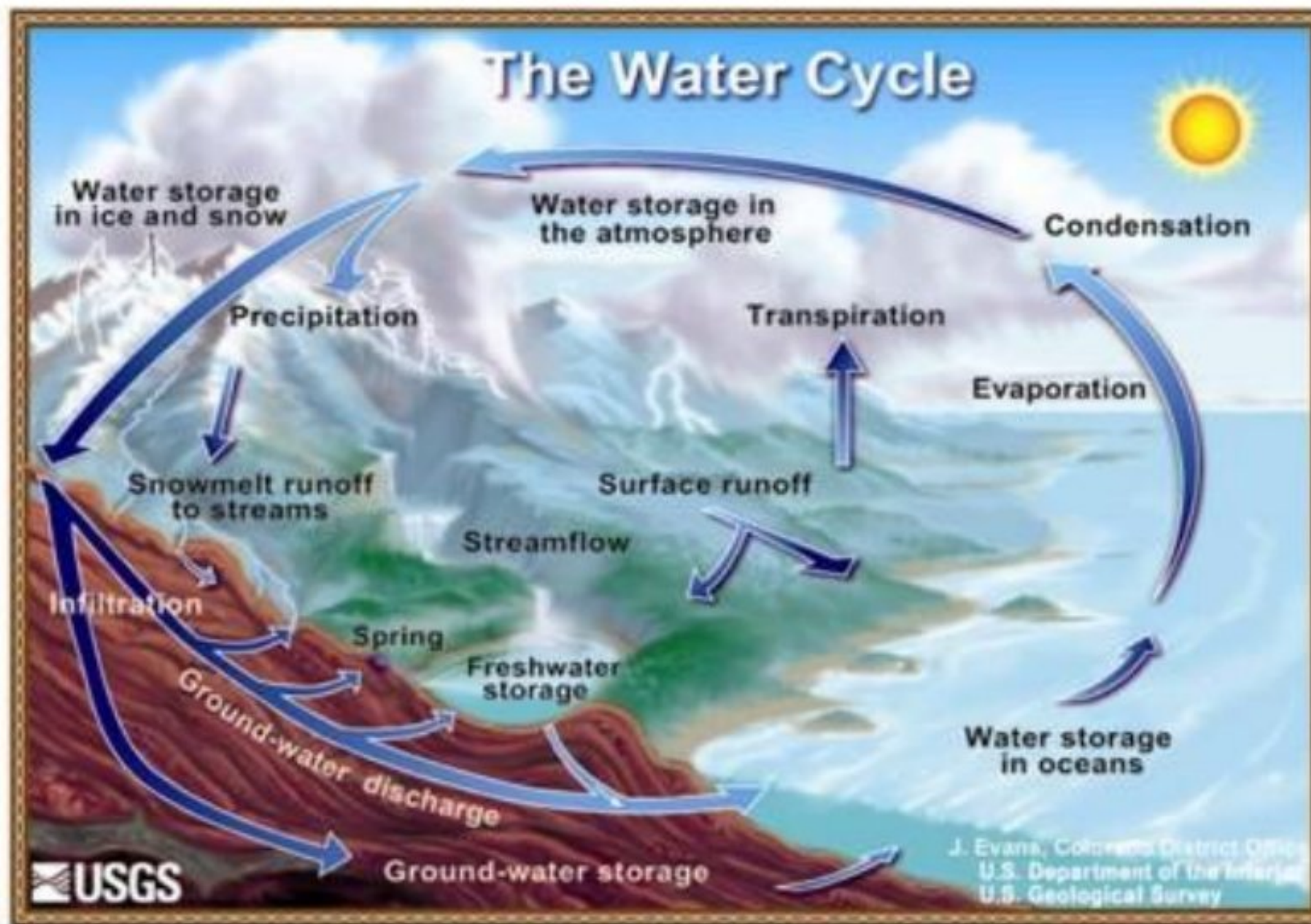
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## Water Cycle – Hydrological cycle

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- There is a continuous exchange of water between the atmosphere, the oceans and the continents through the processes of **evaporation, transpiration, condensation and precipitation**.
- The moisture in the atmosphere is derived from water bodies through **evaporation** and from plants through **transpiration (evapotranspiration)**.
- Evaporated water undergoes **condensation** and forms clouds.
- When saturation is reached, clouds give away water in the form of precipitation.
- Since the total amount of moisture in the entire system remains constant, a balance is required between evapotranspiration and precipitation. The hydrological cycle maintains this balance.



## Water Vapour in Atmosphere

- Water vapour in air varies from **zero to four per cent** by volume of the atmosphere (averaging around **2%** in the atmosphere). Amount of water vapour (Humidity) is measured by, an instrument called **Hygrometer**.

## Significance of Atmospheric Moisture

1. Water vapour absorbs radiation—**both incoming and terrestrial**. It thus plays a crucial role in the **earth's heat budget**.
2. The amount of water vapour present decides the **quantity of latent energy stored** up in the atmosphere for development of storms and cyclones.
3. The atmospheric moisture affects the human body's rate of cooling by influencing the sensible temperature.

## Evaporation

- The oceans covering 71% of the earth's surface hold 97% of all the earth's water reserves.
- Evapotranspiration may be taken as the starting point in the hydrological cycle. The oceans contribute **84%** of the annual total and the continents **16%**.

- The highest annual evaporation occur in the **sub-tropics of the western North Atlantic** and **North Pacific** because of the influence of the **Gulf Stream** and the **Kurishino Current**, and in the **trade wind zone of the southern oceans**.
- The land maximum occurs in equatorial region because of **high insolation** and luxuriant

## Humidity

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- Water vapour present in the air is known as Humidity.

## Absolute Humidity

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- The **actual** amount of the water vapour present in the atmosphere is known as the **absolute humidity**.
- It is the **weight** of water vapour per unit volume of air and is expressed in terms of grams per cubic metre.
- The absolute humidity **differs** from place to place on the surface of the earth.
- The ability of the air to hold water vapour depends entirely on its temperature (**Warm air can hold more moisture than cold air**).

## Relative Humidity

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- The **percentage of moisture present in the atmosphere as compared to its full capacity at a given temperature** is known as the relative humidity.

*Relative Humidity = [Actual amount of water vapor in air (absolute humidity)/humidity at saturation point (the maximum water vapor air can hold at a given temperature)] X 100*

- With the change of air temperature, the capacity to retain moisture increases or decreases and the relative humidity is also affected.
- **Relative humidity is greater over the oceans and least over the continents** (absolute humidity is greater over oceans because of greater availability of water for evaporation).
- The **relative humidity** determines the **amount and rate of evaporation** and hence it is an **important climatic factor**.
- Air containing moisture to its full capacity at a given temperature is said to be '**saturated**'. At this temperature, the air **cannot** hold any additional amount of moisture. Thus, relative humidity of the saturated air is **100%**.

- If the air has half the amount of moisture that it can carry, then it is unsaturated and its relative humidity is only **50%**.

**Relative humidity can be changed in either of the two ways—**

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1. **By adding moisture through evaporation (by increasing absolute humidity):** if moisture is added by evaporation, the relative humidity will increase and vice versa.
2. **By changing temperature of air (by changing the saturation point):** a decrease in temperature (hence, decrease in moisture-holding capacity/decrease in saturation point) will cause an increase in relative humidity and vice versa.

Consider  $1 \text{ m}^3$  of air at a temperature 'T'.

Let us assume that saturation occurs when 0.5 kg of water vapor is present in  $1 \text{ m}^3$  of air.

That is, relative humidity will be 100% if  $1 \text{ m}^3$  of air contains 0.5 kg of water vapor at temperature T (saturation temperature or saturation point).

Assume that  $1 \text{ m}^3$  of air at a given time consists of 0.2 kg of water vapor at a temperature 'T'.

Now the relative humidity = 40 %  $\implies$  0.2 kg of water vapor per  $1 \text{ m}^3$  of air  $\implies$  the air can still hold 0.3 kg of water vapor since saturation occurs at 0.5 kg.

Here,

**Absolute Humidity =  $0.2 \text{ kg/ m}^3$  and**

**Relative Humidity = 40 %**

So, relative humidity is expressed as % whereas absolute humidity is expressed in absolute terms.

Now to make the air saturated (100 % relative humidity),

1. we can add that additional 0.3 kg of water vapor by evaporation. **OR**
2. we can decrease the temperature.

If we decrease the temperature, the saturation point will come down.

Let us assume that the temperature of  $1 \text{ m}^3$  of air is decreased by  $2 \text{ }^\circ\text{C}$ . The water holding capacity will fall due to decrease in temperature. Let us assume that the water holding capacity decreases by  $0.1 \text{ kg per m}^3$  of air per  $1 \text{ }^\circ\text{C}$  fall in temperature.

So, for  $2 \text{ }^\circ\text{C}$  fall in temperature, the fall in water holding capacity is  $0.2 \text{ kg/m}^3$  of air ( $0.1 \text{ kg/m}^3 \times 2$ ).

Hence the new saturation point occurs at  $0.3 \text{ kg/m}^3$  of air [ $0.5 \text{ kg/m}^3 - 0.2 \text{ kg/m}^3$ ].

That is, the 'new saturation point' (relative humidity = 100%) occurs when the water vapor content is  $0.3 \text{ kg per } 1 \text{ m}^3$  of air.

So now we can saturate  $1 \text{ m}^3$  of air by adding just  $0.1 \text{ kg}$  instead of  $0.3 \text{ kg}$  as in the earlier case.

[because, initially, we assumed that  $1 \text{ m}^3$  of air at a given time consists of  $0.2 \text{ kg}$  of water vapor at a temperature 'T'.]

## Dew point

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- The air containing moisture to its full capacity at a given temperature is said to be saturated.
- It means that the air at the given temperature is incapable of holding any additional amount of moisture at that stage.
- The temperature at which saturation occurs in a given sample of air is known as **dew point**.
- **Dew point occurs when Relative Humidity = 100%.**

## Specific Humidity

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- It is expressed as the **weight of water vapour per unit weight of air**.
- Since it is measured in units of weight (usually grams per kilogram), the specific humidity is **not affected by changes in pressure or temperature**.

**Absolute Humidity and Relative Humidity are Variable whereas Specific Humidity is a constant.**